

## Physiological impacts on alkenone paleothermometry

F. G. Prahl

College of Oceanic and Atmospheric Sciences, Oregon State University, Corvallis, Oregon, USA

G. V. Wolfe

Department of Biological Sciences, California State University, Chico, California, USA

M. A. Sparrow

College of Oceanic and Atmospheric Sciences, Oregon State University, Corvallis, Oregon, USA

Received 6 May 2002; revised 17 October 2002; accepted 23 December 2002; published 9 April 2003.

[1] We conducted isothermal (15°C) batch culture experiments with the coccolithophorid *Emiliania huxleyi* (strain NEPCC 55a) to evaluate the extent to which nutrient and light stress contribute to variability in the alkenone unsaturation index  $U_{37}^{K'}$ . Alkenone content and composition were constant throughout exponential growth in both experiments when nutrients (nitrate and orthophosphate) were replete. Stationary phase (nutrient-starved) cells continued to produce alkenones, amassing concentrations ( $\Sigma\text{Alk}$ )  $\geq 3$  times higher than those dividing exponentially (1.5–2 pg cell<sup>-1</sup>), and the  $U_{37}^{K'}$  of “excess” alkenone dropped by 0.11 units. In contrast, 5 days of continuous darkness resulted in a 75% decrease in cellular  $\Sigma\text{Alk}$  and a significant  $U_{37}^{K'}$  increase (+0.11 units). Given an established 0.034 unit/°C response for exponentially growing cells of this strain, the observed range of  $U_{37}^{K'}$  variability at 15°C corresponds to an uncertainty of  $\pm 3.2^\circ\text{C}$  in predicted growth temperature. This level of variability matches that of the global  $U_{37}^{K'}$  annual mean sea surface temperature calibration for surface marine sediments, begging the question: What is the physiological condition of alkenone-producing cells exported to marine sediments? Comparison of our laboratory results for a strain of *E. huxleyi* isolated from the subarctic Pacific Ocean with depth profiles for alkenones in surface waters from two contrasting sites in the northeast Pacific Ocean suggests that the answer to this question depends on the ocean regime considered, a possibility with significant bearing on how stratigraphic  $U_{37}^{K'}$  records in marine sediments are to be interpreted paleoceanographically. **INDEX TERMS:** 4267 Oceanography: General: Paleoclimatology; 4808 Oceanography: Biological and Chemical: Chemical tracers; 4850 Oceanography: Biological and Chemical: Organic marine chemistry; 4855 Oceanography: Biological and Chemical: Plankton; 4863 Oceanography: Biological and Chemical: Sedimentation

**Citation:** Prahl, F. G., G. V. Wolfe, and M. A. Sparrow, Physiological impacts on alkenone paleothermometry, *Paleoceanography*, 18(2), 1025, doi:10.1029/2002PA000803, 2003.

### 1. Introduction

[2] Unsaturation ratios ( $U_{37}^{K'}$ ) in long-chain alkenones from haptophyte algae serve stratigraphically in marine sediments as a paleorecord for sea surface temperature (SST) [Brassell *et al.*, 1986]. Validated globally by a strong statistical correlation between  $U_{37}^{K'}$  measures in surface sediments and overlying annual mean SST [Muller *et al.*, 1998], this method has now been employed to yield paleotemperature records for the Holocene [e.g., Kennedy and Brassell, 1992], Pleistocene [e.g., Bard *et al.*, 1997; Lyle *et al.*, 1992; Rostek *et al.*, 1993; Sachs and Lehman, 1999], and Miocene [e.g., Pagani *et al.*, 1999] periods. However, the statistical variability of the  $U_{37}^{K'}$ -SST calibration adds uncertainty to its paleothermometric accuracy [Prahl *et al.*, 2000]. A combination of environmental and biological factors almost certainly contributes at some level to this variability. Although the  $U_{37}^{K'}$  value appears resistant to diagenetic change [Grimm *et al.*, 2000], not all algae capable of alkenone biosyn-

thesis [Brassell, 1993] and of potentially contributing a signal to sediments display identical temperature responses for  $U_{37}^{K'}$  [Conte *et al.*, 1998]. Also, in addition to these apparent genetic differences, physiological concerns within a given genotype are also now evident [Conte *et al.*, 1998; Epstein *et al.*, 1998].

[3] Here we present results from two types of batch culture experiments with a specific strain of major alkenone producer, *Emiliania huxleyi*, whose calibration curve represents global data remarkably well [Muller *et al.*, 1998]. These results show how physiological factors such as nutrient and light availability impact  $U_{37}^{K'}$  ratios recorded in growing and stationary phase cells. Our laboratory-controlled experiments were designed to illuminate the first-order impact of physiology on biomarker synthesis by alkenone-producing haptophytes in natural waters and the integrity of thermometric interpretations based on the analysis of alkenone unsaturation patterns ( $U_{37}^{K'}$ ) in sedimentary samples. Nutrient stress like that imposed on cells by our first experiment can be encountered at the end of batch-culture-like algal blooms stimulated by the upwelling of fertile waters along coastal

margins. Light-limited growth conditions, potentially as extreme as those posed by the second experiment, are encountered by algal cells residing in deep chlorophyll maxima prevalent throughout the open ocean [e.g., *Longhurst and Harrison*, 1989] or transported to depth by deep mixing events along margins [e.g., *Murphy and Cowles*, 1997]. Although our findings by no means undermine the utility and value of alkenone paleothermometry, they do highlight a clear need to carefully consider such nonthermal environmental variables when stratigraphic  $U_{37}^{K'}$  profiles from different sedimentary regimes are specifically interpreted, particularly in context with other paleoceanographic proxies [*Mix et al.*, 2000].

## 2. Methods

### 2.1. Culture Conditions

[4] NEPCC 55a, the subarctic Pacific ( $50^{\circ}11'N$ ,  $144^{\circ}57'W$ ) strain of *Emiliania huxleyi* used to establish the commonly employed  $U_{37}^{K'}$ -growth temperature calibration [*Prahl and Wakeham*, 1987], was cultured in 500 mL Pyrex flasks in a constant temperature water bath ( $15^{\circ}C$ ) using Cool-White fluorescent lights ( $\sim 50 \mu E m^{-2} sec^{-1}$ ; 12:12 daily light cycle) and a standard growth media derivative (f/20-Si, see [http://ccmp.bigelow.org/CI/CI\\_01e.html](http://ccmp.bigelow.org/CI/CI_01e.html)). NEPCC 55a was obtained as CCMP1742 from the Provasoli-Guillard National Center for Culture of Marine Phytoplankton. Culture flasks were stirred by gentle swirling at the same time each day and immediately subsampled using a sterile pipette for daily determination of cell density, nitrate and orthophosphate concentrations, cellular organic carbon content, and cellular alkenone content and composition.

[5] At each time point, cell density was determined microscopically using a haemocytometer. Cells in known volumes of culture were also filtered using gentle vacuum onto two precombusted ( $450^{\circ}C$ , 8 hour), 25 mm diameter glass fiber filters (GFF). Each GFF was wrapped in aluminum foil and stored frozen ( $-80^{\circ}C$ ) until needed for organic chemical analysis. A volume of filtrate from each sample was also collected in an acid-cleaned, 25-mL polyethylene vial and stored frozen until needed for nutrient analysis.

### 2.2. Chemical Analyses

[6] One GFF was fumed overnight with concentrated hydrochloric acid to remove any trace of inorganic carbon and was then analyzed for organic carbon content using a standard high-temperature combustion method [*Verardo et al.*, 1990]. The other filter was extracted ultrasonically three successive times using a 75:25 mixture of methylene chloride and methanol (20 mL each). Upon the addition of water (10 mL), combined extracts (60 mL total) were partitioned into hexane. After drying over granular, anhydrous sodium sulfate, the hexane solution was then concentrated by rotary evaporation and transferred into a solvent-cleaned, 1 dram borosilicate glass vial with a Teflon-lined cap. Upon evaporation of the solvent under a gentle stream of prepurified  $N_2$  the resultant total lipid residues were derivatized using bis-trifluorotrimethylacetamide with 1% trimethylchlorosilane (Sigma Chemicals) and

analyzed for alkenone content and composition using capillary gas chromatography with flame ionization detection (GC-FID). All reported alkenone concentrations have been corrected for recovery efficiency (typically 80–90%) of an internal standard (nonadecan-2-one) added to each sample prior to extraction. Further details of the alkenone work-up and GC-FID analysis procedure are presented by *Prahl et al.* [1988]. Frozen filtrate from each sampling time point was thawed and analyzed for nitrate and orthophosphate concentration by routine, autoanalyzer-based, colorimetric methods [*Strickland and Parsons*, 1972].

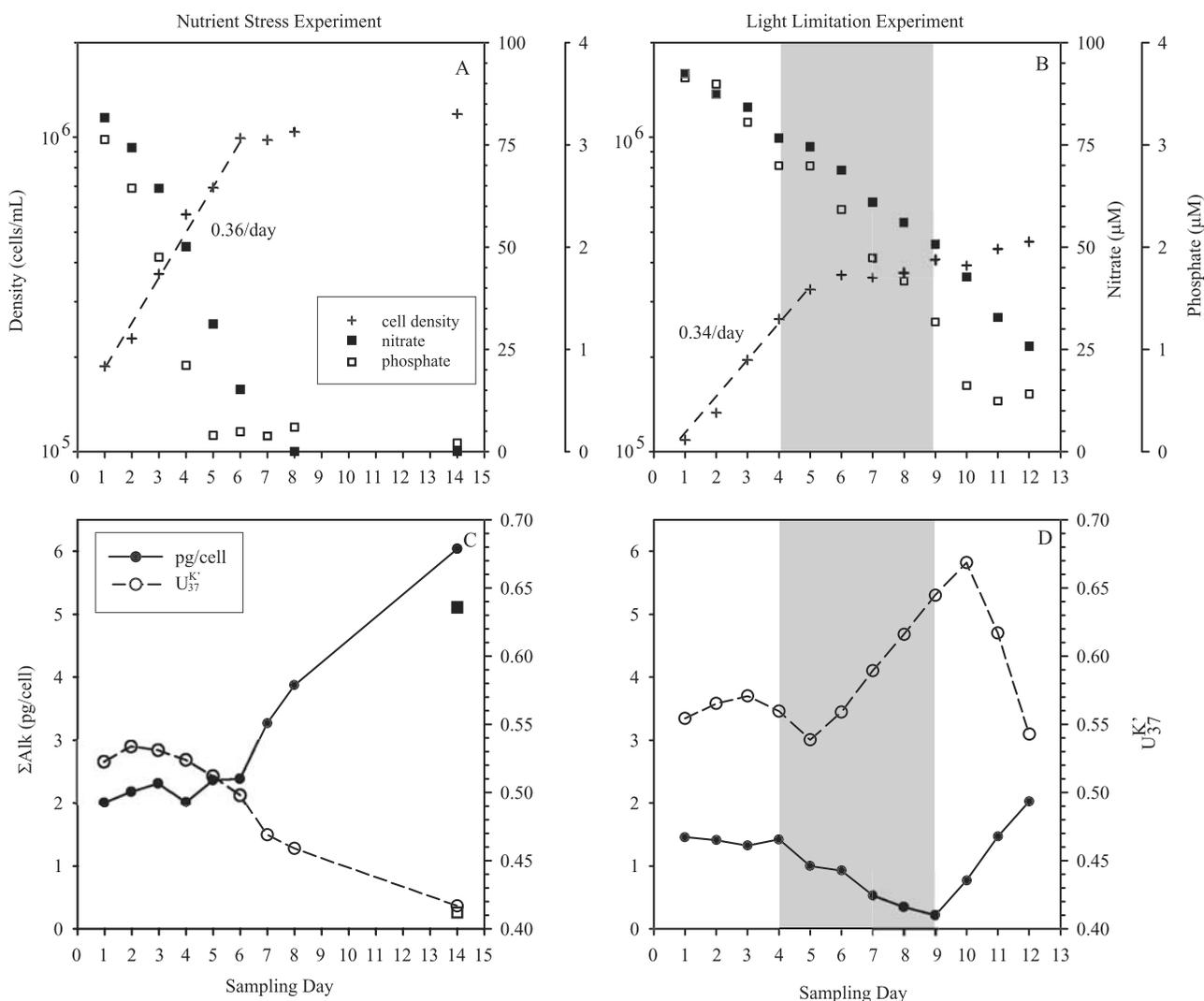
## 3. Results

[7] Two different types of isothermal ( $15^{\circ}C$ ) batch culture experiments were conducted with NEPCC 55a. Alkenone content and composition were tracked (1) in cells dividing exponentially until growth was limited by nutrient depletion (nutrient stress experiment) and (2) in cells shifted at midexponential growth to prolonged, continuous darkness (light limitation experiment).

[8] Cells at the start of both experiments grew exponentially at a rate of  $\sim 0.35$  divisions  $d^{-1}$  (Figures 1a and 1b). In the nutrient stress experiment, cells maintained this growth rate until they stripped the medium of its limiting nutrient, orthophosphate. Cells stopped dividing 1 day after orthophosphate was depleted but continued to consume nitrate for another day until it was also exhausted (Figure 1a). The molar ratio for nitrate to orthophosphate drawdown during the exponential growth period was 15.1:1, essentially Redfield stoichiometry.

[9] In the light limitation experiment the growth rate of  $\sim 0.35$  divisions  $d^{-1}$  was maintained until the cells were placed in continuous darkness, upon which cell division ceased after the first day (Figure 1b). When cells were reexposed after 5 days to their prior daily light-dark cycle, they resumed growth but at  $<50\%$  of the initial rate. Nitrate and orthophosphate uptake was evident in nondividing cells throughout the darkness phase of the experiment (Figure 1b). Nutrient uptake in darkness occurred in the molar proportion of  $\sim 16.6:1$ , only slightly higher than in cells growing exponentially under lighted conditions.

[10] In both experiments, total alkenone concentrations ( $\Sigma Alk$ ,  $pg cell^{-1}$ ) were relatively constant in exponentially growing cells, but absolute concentrations differed between experiments, with those in the nutrient stress experiment containing  $\sim 40\%$  more (compare Figures 1c and 1d). Upon nutrient depletion, cellular alkenone concentration increased steadily by a factor of  $\sim 3$  over an 8-day period (Figure 1c). This increase was accompanied by a systematic decline in the alkenone unsaturation index  $U_{37}^{K'}$  ( $[K37:2]/([K37:2] + [K37:3])$ ), a trend driven by selective accumulation of the tri-unsaturated (K37:3) relative to the di-unsaturated (K37:2)  $C_{37}$  methyl ketone (Table 1). Upon exposure to continuous darkness, cellular alkenone concentration decreased steadily by  $\sim 80\%$  over the 5-day period (Figure 1d). This decrease was accompanied by a pronounced increase in  $U_{37}^{K'}$ , a trend driven by selective loss of K37:3 (Table 1). Upon reinstating a daily light-dark cycle, cells rapidly accumulated alkenones and, after 3 days, displayed a concentration and  $U_{37}^{K'}$  like that observed



**Figure 1.** Batch culture experiments conducted with *E. huxleyi* strain NEPCC 55a to test the effect of (a, c) nutrient depletion (nutrient stress experiment) and (b, d) darkness (light limitation experiment) on cellular alkenone content and composition. The shaded areas in Figures 1b and 1d indicate the 5-day period in the light limitation experiment when the culture flask was covered with sufficient aluminum foil to exclude all light but not to impede gaseous exchange between the culture media and the atmosphere. Figures 1a and 1b display growth curves for each experiment. Dashed lines identify the portion of each curve used to calculate exponential growth rates ( $\sim 0.35$  divisions  $d^{-1}$ ). Data depicting the drawdown of nitrate and orthophosphate concentration in each experiment are also plotted. Figures 1c and 1d illustrate how total  $C_{37-39}$  alkenone concentration ( $\Sigma Alk$ ,  $pg\ cell^{-1}$ ) and values for the unsaturation index,  $U_{37}^{K'}$ , evolved during the course of each experiment. The solid and open squares in Figure 1c depict  $C_{37-39}$  alkenone concentration and  $U_{37}^{K'}$ , respectively, measured in stationary phase cells from day 14 of the nutrient stress experiment that were subsequently put into continuous darkness for 5 days prior to sampling for analysis.

in exponentially dividing cells at the start of the experiment (Figure 1d).

[11] The two types of experiments described above have been repeated now several times using culture medium with different starting proportions of nitrate to orthophosphate (F. G. Prahl et al., unpublished data, 2001, 2002). The salient features of these experiments are the same as what has just been described. Detailed results from the complete set of

experiments will be discussed more specifically in a future manuscript.

#### 4. Discussion

[12] Results from our batch culture experiments add to the growing body of evidence that nutrient or light stress can affect cellular burdens and isomeric composition of long-

**Table 1.** Summary of Results From the Nutrient Stress and Light Limitation Experiments Conducted With Strain NEPCC 55a of *E. huxleyi* Grown in Laboratory-Controlled Batch Cultures

Experiment	Sampling Day	POC, $\text{pg cell}^{-1}$	$\Sigma\text{Alk}^a$ , % of C	K37:2, $\text{pg cell}$	K37:3, $\text{pg cell}$	K37, <sup>b</sup> %	K38m, <sup>b</sup> %	K38e, <sup>b</sup> %	K39, <sup>b</sup> %
Nutrient stress	1	9.6	16	0.607	0.664	64	13	20	2.6
Nutrient stress	2	10.1	17	0.660	0.754	64	13	20	2.2
Nutrient stress	3	8.9	21	0.675	0.763	62	14	22	2.0
Nutrient stress	4	8.1	20	0.625	0.686	65	12	21	2.1
Nutrient stress	5	9.6	19	0.748	0.784	65	12	21	2.2
Nutrient stress	6	9.4	20	0.764	0.758	63	12	22	2.4
Nutrient stress	7	11.9	22	1.06	0.942	61	13	23	2.8
Nutrient stress	8	12.8	24	1.26	1.07	60	14	24	3.3
Nutrient stress	14	12.1	39	1.98	1.42	56	15	26	4.2
Nutrient stress (5 days darkness)	19	8.9	46	1.63	1.14	53	15	28	4.3
Light limitation (light)	1	7.3	15	0.392	0.488	62	12	20	5.7
Light limitation (light)	2	6.4	17	0.412	0.535	68	11	20	1.6
Light limitation (light)	3	6.2	17	0.378	0.502	67	11	20	1.9
Light limitation (light to dark)	4	7.7	15	0.413	0.526	67	11	20	2.0
Light limitation (dark)	5	5.9	13	0.298	0.348	65	12	21	1.7
Light limitation (dark)	6	5.6	13	0.242	0.306	61	12	23	3.1
Light limitation (dark)	7	6.3	7	0.122	0.174	56	14	28	2.4
Light limitation (dark)	8	6.5	4	0.068	0.109	51	14	32	3.1
Light limitation (dark to light)	9	5.3	4	0.039	0.071	45	14	35	5.6
Light limitation (light)	10	8.0	8	0.141	0.284	55	15	25	4.3
Light limitation (light)	11	8.9	13	0.330	0.531	59	15	23	3.2
Light limitation (light)	12	9.6	17	0.536	0.636	58	15	23	3.7

<sup>a</sup> $\Sigma\text{Alk}$   $\equiv$  total concentration of  $C_{37}$ ,  $C_{38}$ , and  $C_{39}$  methyl and ethyl ketones expressed as a percentage of total cellular organic carbon; calculation made assuming the mass of each compound is  $\sim 80\%$  carbon by weight.

<sup>b</sup>%K37, %K38m, %K38e, %K39  $\equiv$  concentration of  $C_{37}$  methyl,  $C_{38}$  methyl,  $C_{38}$  ethyl, and  $C_{39}$  ethyl ketones as a percentage of total alkenone concentration, respectively.

chain alkenones in *E. huxleyi*. Buildup of alkenone concentration in nutrient-stressed cells has been reported previously in other strains [Conte *et al.*, 1998; Epstein *et al.*, 1998]. Significant change in  $U_{37}^{K'}$  has also been noted during this concentration buildup, although the shift reported by Epstein *et al.* [1998] was opposite in direction to what Conte *et al.* [1998] documented and our work now shows. The cause for the apparent biosynthetic difference between strains is unclear and needs further investigation.

[13] Recently, loss of alkenone concentration in cells kept in continuous darkness has also been reported for NEPCC 55a and another *E. huxleyi* strain [Epstein *et al.*, 2001]. Unlike the case for our light limitation experiment, however, Epstein *et al.* focused on cells that had been in stationary growth for at least a week and had no nitrate or orthophosphate left in their medium. The magnitude of loss of cellular alkenone concentration was comparable to what we observed (i.e.,  $\sim 80\%$ ) but occurred over a period of time 4 times longer. Although selective loss of K37:3 was observed, the degree of selectivity was much less, resulting in a  $U_{37}^{K'}$  shift of only 0.04 units. In our nutrient stress experiment with NEPCC 55a we reexamined the effect of darkness on alkenone content and composition of stationary phase cells. We found that the contrasts just described are reproducible. Stationary phase cells from day 14 that were kept for 5 additional days in total darkness prior to analysis showed a more minor loss of alkenone content ( $\sim 15\%$ ) with negligible impact on  $U_{37}^{K'}$  (Figure 1c, compare corresponding solid and open symbols).

[14] We also evaluated how alkenone contribution to total cellular carbon ( $\Sigma\text{Alk}$  as percent of C, Table 1) varied as a function of nutrient stress and light limitation. As documented in prior work with *E. huxleyi* [Conte *et al.*, 1998;

Prahl *et al.*, 1988], alkenones comprised a significant fraction of total cellular carbon in exponentially growing cells (i.e., 15–20%, Table 1). Under nutrient stress, when stationary phase cells concentrated even more of these lipids (Figure 1c), their contribution to total cellular carbon more than doubled, reaching 39% in the final day of our first experiment (Table 1). Under light deprivation, when non-dividing, nutrient replete cells appear to consume these biochemicals (Figure 1d), their contribution to total cellular carbon is reduced significantly, reaching 4% on the fifth day of darkness in our second experiment (Table 1). Notably, alkenone contribution to total cellular carbon was not reduced, but rather increased somewhat, in the case where nutrient-limited, stationary phase cells from day 14 of our first experiment were subjected to darkness for 5 additional days prior to sampling and analysis (Table 1). For physiological reasons yet to be specifically determined, the effect of darkness on alkenone content and composition appears greatest for nutrient replete, exponentially growing cells.

[15] Alkenones are typically biosynthesized by *E. huxleyi* and selected other haptophytes as a set of  $C_{37}$ – $C_{38}$  methyl and  $C_{38}$ – $C_{39}$  ethyl ketones [Brassell, 1993]. In order to illuminate the effect of physiological factors such as nutrient stress and light limitation on alkenone biosynthesis we formulated four compositional indices, i.e., %K37, %K38m, %K38e, and %K39. The compositional index %K37 is defined by the relative abundance of  $C_{37}$  to total  $C_{37}$ ,  $C_{38}$ , and  $C_{39}$  alkenones. The indices %K38m, %K38e, and %K39 are defined accordingly for the  $C_{38}$  methyl,  $C_{38}$  ethyl, and  $C_{39}$  alkenones, respectively. An examination of the data in Table 1 shows that the pronounced alkenone concentration increase in illuminated, nutrient-limited, stationary phase cells and its decrease in light-deprived,

nutrient replete stationary phase cells has a similar consequence on %K37: a decrease in its value. The cause for the decrease is not the same for the two cell types, however: In the case of nutrient stress the decrease results from selective biosynthetic production of  $C_{38}$  methyl/ethyl and  $C_{39}$  ethyl ketones, while in the case of light limitation the decrease results from selective metabolic removal of the  $C_{37}$  methyl ketones (Table 1).

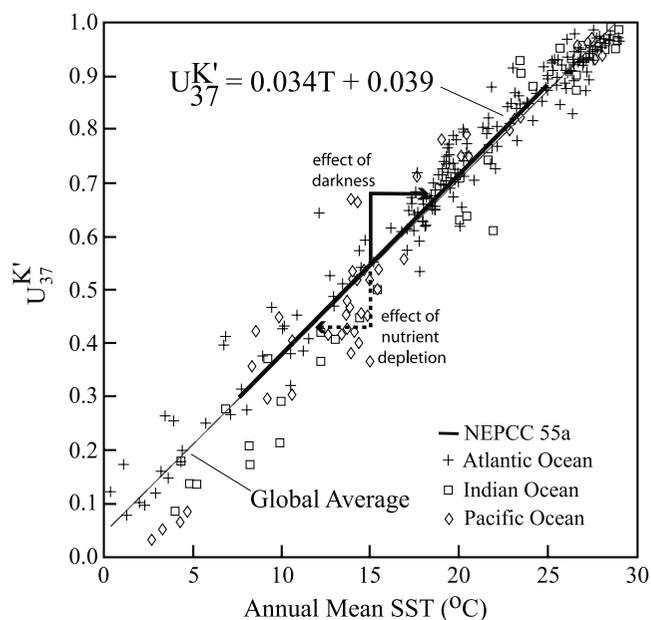
#### 4.1. Physiological Function of Alkenones

[16] The observations from our batch culture experiments are consistent with the suggested role of alkenones as cellular energy stores [Epstein *et al.*, 2001; Pond and Harris, 1996]. *E. huxleyi* does not biosynthesize significant levels of triacylglycerides (TAG) [Volkman *et al.*, 1986], the neutral lipids that other algal phyla typically accumulate when cells are nutrient stressed [Lombardi and Wangersky, 1991; Parrish and Wangersky, 1987]. On this basis, Epstein *et al.* [2001] speculated that alkenones in *E. huxleyi* are biochemical surrogates for TAG. However, the analogy may not be exact, as diatoms accumulate TAG as conspicuous droplets in transmission electron microscopy (TEM) images [Sicko-Goad *et al.*, 1988]. Review of TEM literature revealed no report of such features in *E. huxleyi* images [Pienaar, 1994; van der Wal *et al.*, 1985]. The apparent absence of such a feature is quite noteworthy given that  $C_{37-39}$  alkenones are major contributors to cellular carbon in these organisms, comprising 10–20% of that in exponentially growing cells [Conte *et al.*, 1998; Prahl *et al.*, 1988] and up to 40% of that in nutrient-stressed, stationary phase cells (Table 1).

[17] On the basis of the response of alkenone unsaturation to temperature, Brassell *et al.* [1986] originally conjectured that these biomarkers serve a structural role in the cell by helping in some way to regulate membrane fluidity [Harwood and Russell, 1984]. Our results argue that alkenones are not used strictly for purposes of cellular architecture and biophysical maintenance of membrane properties. Nonetheless, they also do not discount the possibility that these compounds are associated in some way with a particular membrane in cellular substructure (e.g., chloroplasts). Clearly, more laboratory-controlled study is warranted to elucidate specifically how alkenones are contained and how they function physiologically within the cells of the select group of haptophytes capable of this biosynthesis [Brassell, 1993].

#### 4.2. Oceanographic Evidence for Physiological Uncertainty in Paleothermometer Calibration

[18] Despite well-documented genetic variation in  $U_{37}^{K'}$ -T calibration among *E. huxleyi* strains [Conte *et al.*, 1998], a remarkable linear correlation exists between nearly global  $U_{37}^{K'}$  measures in marine surface sediments and annual mean SST in overlying waters that is statistically indistinguishable from the  $U_{37}^{K'}$ -growth temperature calibration for NEPCC 55a [Muller *et al.*, 1998] (Figure 2). Our data support prior studies [Conte *et al.*, 1998; Epstein *et al.*, 1998] which suggest that physiological factors such as nutrient and light availability could contribute to the apparent variability in the core top data set. Although maintained isothermally, the observed 0.11 unit change in each

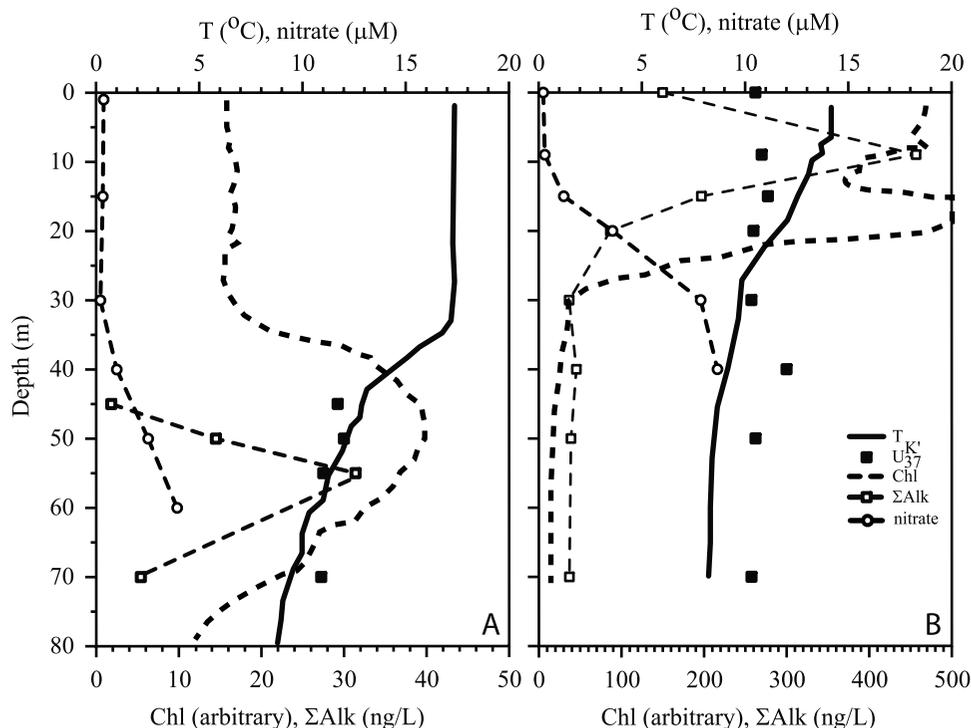


**Figure 2.** Graph displaying the complete data set used for the global core top calibration of  $U_{37}^{K'}$  versus annual mean SST [Muller *et al.*, 1998]. Comparison shows that the best linear fit of these data (shaded dashed line, “global average” defined by  $U_{37}^{K'} = 0.033T + 0.069$  from Muller *et al.* [1998]) is virtually indistinguishable from the original batch culture calibration of *E. huxleyi* strain NEPCC 55a (bold line,  $U_{37}^{K'} = 0.034T + 0.039$  from Prahl *et al.* [1988]). The bold and dotted arrows depict quantitative effects of nutrient depletion and darkness on  $U_{37}^{K'}$ , respectively, documented in the present study by our isothermal batch culture experiments with NEPCC 55a.

of our experiments corresponds to a perceived  $\sim 3^{\circ}\text{C}$  change in growth temperature if translated using the established calibration for NEPCC 55a ( $U_{37}^{K'} = 0.034T + 0.039$  [Prahl *et al.*, 1988]). Superimposing the similar magnitude but opposite quantitative effects of nutrient stress and light limitation on  $U_{37}^{K'}$  documented by our experiments on the plot in Figure 2 shows that potentially all of the variability in the core top calibration could be accounted for by algal physiology.

[19] This prospect begs an answer to the following biological oceanographic question: What is the physiological condition of cells exporting alkenones to marine sediments? The answer to this question is most relevant to paleoceanographic temperature reconstructions based on  $U_{37}^{K'}$  [Conte *et al.*, 1998] and may depend on what particular ocean location is under consideration [e.g., Conte *et al.*, 2001]. Comparison of depth profiles for alkenone content and composition in suspended particulate material from surface waters collected at two different northeast Pacific oceanographic sites provides support, albeit yet circumstantial, for a geographically dependent answer to this question.

[20] Open ocean waters in the temperate northeast Pacific ( $\sim 40^{\circ}\text{N}$ ,  $132^{\circ}\text{W}$ ) are seasonally stratified, and in summer through early fall they display a subsurface chlorophyll maximum (SCM) at the top of a nutricline within the upper



**Figure 3.** Depth profiles for temperature (solid line), nitrate concentration (open circles), chlorophyll fluorescence (Chl) (dashed line), and total  $C_{37-39}$  alkenone concentration ( $\Sigma$ Alk) (open squares) in suspended particulate materials (SPM), measured in the late 1980s in surface waters at two sites in the northeast Pacific Ocean. (a) Temperate gyre location ( $\sim 40^{\circ}\text{N}$ ,  $132^{\circ}\text{W}$ ) [Prahl *et al.*, 2001] and (b) continental margin location ( $46^{\circ}45'\text{N}$ ,  $125^{\circ}5'\text{W}$ ) off the coast of Washington State (F. G. Prahl *et al.*, unpublished data, 2001, 2002). Solid squares depict growth temperature estimates for alkenone producers at the depth of SPM collection using the established  $U_{37}^{K'}$  calibration for NEPCC 55a (i.e.,  $U_{37}^{K'} = 0.034T + 0.039$  [Prahl *et al.*, 1988]).

thermocline. In September 1989 we found measurable alkenone concentrations at this site only in the SCM, with  $U_{37}^{K'}$ -based growth temperature estimates in good agreement with actual water temperature measured in this feature (Figure 3a). In contrast, ocean waters along the northeast Pacific continental margin experience seasonal upwelling of nutrients [Hickey, 1979] and consequently, episodic phytoplankton blooms throughout summer and early fall [Landry and Hickey, 1989]. At one such site off the Washington coast ( $46^{\circ}45'\text{N}$ ,  $125^{\circ}5'\text{W}$ ) in August 1988 we found high alkenone concentrations in near-surface waters containing very high chlorophyll and low nitrate concentrations (Figure 3b).  $U_{37}^{K'}$ -based growth temperature estimates were  $\sim 3^{\circ}\text{C}$  lower than actual surface water temperatures. We have also made similar observations at continental margin sites in the northeast Pacific off northern California ( $\sim 40^{\circ}\text{N}$ ) (F. G. Prahl *et al.*, unpublished data, 1989). When these observations are considered in context with findings from our batch culture experiments, alkenone producers at the open ocean site in the temperate northeast Pacific would seem to have a supply of nutrient and light sufficient for perhaps slow but steady state, continuous culture-like growth [e.g., Popp *et al.*, 1998] and thereby export a  $U_{37}^{K'}$  signal to the seafloor that predominantly records water temperature at the depth of production. Alkenone signatures in continental margin waters of the northeast Pacific, on the

other hand, seem consistent with that of cells experiencing significant nutrient stress under episodic, batch-culture-like growth conditions (e.g., Figures 1a and 1c). If so, the  $U_{37}^{K'}$  signal exported to the seafloor at such sites quite significantly underestimates actual growth temperature at the depth of production (i.e.,  $\sim 3^{\circ}\text{C}$ ). Recognizing whether the bias in the temperature assessment is systematic or variable through time would then pose a significant challenge to paleoceanographic research.

[21] The physiological interpretation applied to these water column profiles is certainly not yet unequivocal and warrants further evaluation. Nonetheless, it is now defensible based on results from our batch culture experiments and the growing body of biogeochemical literature on these molecules, and it is therefore reasonable to pose. If proven in the future to be correct, it bears significantly on how specifically stratigraphic  $U_{37}^{K'}$  measurements in marine sediments, particularly those accumulating near ocean margins, can be read as a paleorecord for absolute changes in annual mean SST.

## 5. Conclusions

[22] The findings from our work with laboratory batch cultures of a single strain of *E. huxleyi* by no means discount the use of  $U_{37}^{K'}$  as an insightful paleotemperature

proxy. However, they clearly show that the absolute value for  $U_{37}^{K'}$  recorded in sediments may not be set exclusively by growth temperature or by export of a signal from cells necessarily productive at the sea surface. Further research is warranted to refine the understanding of alkenone biosynthesis and its physiological control as well as the understanding of environmental controls on organic matter export from the euphotic zone to sediments. This need for further research is now particularly required if our paleoceanographic

goal [Mix *et al.*, 2000] is to use the alkenone fossil record for more than just a measure of relative changes in annual mean SST through time [Brassell *et al.*, 1986; Muller *et al.*, 1998].

[23] **Acknowledgments.** A collaborative National Science Foundation grant OCE-9986306 to F.G.P. and G.V.W. provided all funding for this research. We are also grateful to Andrew Ross (COAS/OSU) for the nutrient analyses.

## References

- Bard, E., F. Rostek, and C. Sonzogni, Interhemispheric synchrony of the last deglaciation inferred from alkenone palaeothermometry, *Nature*, 385, 707–710, 1997.
- Brassell, S. C., Applications of biomarkers for delineating marine paleoclimatic fluctuations during the Pleistocene, in *Organic Geochemistry: Principles and Applications*, edited by M. H. Engel and S. A. Macko, pp. 699–738, Plenum, New York, 1993.
- Brassell, S. C., G. Eglinton, I. T. Marlowe, U. Pflaumann, and M. Sarthain, Molecular stratigraphy: A new tool for climatic assessment, *Nature*, 320, 129–133, 1986.
- Conte, M. H., A. Thompson, D. Lesley, and R. Harris, Genetic and physiological influences on the alkenone/alkenoate versus growth temperature relationship in *Emiliania huxleyi* and *Gephyrocapsa oceanica*, *Geochim. Cosmochim. Acta*, 62, 51–68, 1998.
- Conte, M. H., J. C. Weber, L. L. King, and S. G. Wakeham, The alkenone temperature signal in western North Atlantic surface waters, *Geochim. Cosmochim. Acta*, 65, 4275–4287, 2001.
- Epstein, B. L., S. D'Hondt, J. G. Quinn, J. Zhang, and P. E. Hargraves, An effect of dissolved nutrient concentrations on alkenone-based temperature estimates, *Paleoceanography*, 13, 122–126, 1998.
- Epstein, B. L., S. D'Hondt, and P. E. Hargraves, The possible metabolic role of  $C_{37}$  alkenones in *Emiliania huxleyi*, *Org. Geochem.*, 32, 867–875, 2001.
- Grimalt, J. O., J. Rullkötter, M.-A. Sicre, R. Summons, J. Farrington, H. R. Harvey, M. Goñi, and K. Sawada, Modifications of the  $C_{37}$  alkenone and alkenoate composition in the water column and sediment: Possible implications for sea surface temperature estimates in paleoceanography, *Geochem. Geophys. Geosyst.*, 1, Paper number 2000GC000053, 2000.
- Harwood, J. L., and N. J. Russell, *Lipids in Plants and Microbes*, 162 pp., Allen and Unwin, Concord, Mass., 1984.
- Hickey, B. M., The California Current System—Hypothesis and facts, *Prog. Oceanogr.*, 8, 191–271, 1979.
- Kennedy, J. A., and S. C. Brassell, Molecular records of twentieth-century El Niño events in laminated sediments from the Santa Barbara basin, *Nature*, 357, 62–64, 1992.
- Landry, M. R., and B. M. Hickey (Eds.), *Coastal Oceanography of Washington and Oregon*, 607 pp., Elsevier Sci., New York, 1989.
- Lombardi, A. T., and P. J. Wangersky, Influence of phosphorus and silicon on lipid class production by the marine diatom *Chaetoceros gracilis* grown in turbidostat cage cultures, *Mar. Ecol. Prog. Ser.*, 77, 39–47, 1991.
- Longhurst, A. R., and W. G. Harrison, The biological pump: Profiles of plankton production and consumption in the upper ocean, *Prog. Oceanogr.*, 22, 47–123, 1989.
- Lyle, M. W., F. G. Prahl, and M. A. Sparrow, Upwelling and productivity changes inferred from a temperature record in the central equatorial Pacific, *Nature*, 355, 812–815, 1992.
- Mix, A. C., E. Bard, G. Eglinton, L. D. Keigwin, A. C. Ravelo, and Y. Rosenthal, Alkenones and multiproxy strategies in paleoceanographic studies, *Geochem. Geophys. Geosyst.*, 1, Paper number 2000GC000056, 2000.
- Muller, P. J., G. Kirst, G. Ruhland, I. von Storch, and A. Rosell-Mele, Calibration of the alkenone paleotemperature index  $U_{37}^{K'}$  based on core-tops from the eastern South Atlantic and the global ocean (60°N–60°S), *Geochim. Cosmochim. Acta*, 62, 1757–1772, 1998.
- Murphy, A. M., and T. J. Cowles, Effects of darkness on multi-excitation in vivo fluorescence and survival in a marine diatom, *Limnol. Oceanogr.*, 42, 1444–1453, 1997.
- Pagani, M., M. A. Arthur, and K. H. Freeman, Miocene evolution of atmospheric carbon dioxide, *Paleoceanography*, 14, 273–292, 1999.
- Parrish, C. C., and P. J. Wangersky, Particulate and dissolved lipid classes in cultures of *Phaeodactylum tricornutum* grown in cage culture turbidostats with a range of nitrogen supply rates, *Mar. Ecol. Prog. Ser.*, 35, 119–128, 1987.
- Pienaar, R. N., Ultrastructure and calcification of coccolithophores, in *Coccolithophores*, edited by A. Winter and W. G. Siesser, pp. 13–37, Cambridge Univ. Press, New York, 1994.
- Pond, W. D., and R. P. Harris, Lipid composition of the coccolithophore *Emiliania huxleyi* and its possible ecophysiological significance, *J. Mar. Biol. Assoc. U. K.*, 76, 579–594, 1996.
- Popp, B. N., F. Kenig, S. G. Wakeham, E. A. Laws, and R. R. Bidigare, Does growth rate affect ketone unsaturation and intracellular carbon isotopic variability in *Emiliania huxleyi*?, *Paleoceanography*, 13, 35–41, 1998.
- Prahl, F. G., and S. G. Wakeham, Calibration of unsaturation patterns in long-chain ketone compositions for palaeotemperature assessment, *Nature*, 330, 367–369, 1987.
- Prahl, F. G., L. A. Muehlhausen, and D. L. Zahnle, Further evaluation of long-chain alkenones as indicators of paleoceanographic conditions, *Geochim. Cosmochim. Acta*, 52, 2303–2310, 1988.
- Prahl, F. G., T. Herbert, S. C. Brassell, N. Ohkouchi, M. Pagani, D. Repeta, A. Rosell-Melé, and E. Sikes, Status of alkenone paleothermometer calibration: Report from Working Group #3, *Geochem. Geophys. Geosyst.*, 1, Paper number 2000GC000058, 2000.
- Prahl, F. G., C. H. Pilskaln, and M. A. Sparrow, Seasonal record for alkenones in sedimentary particles from the Gulf of Maine, *Deep Sea Res. Part I*, 48, 515–528, 2001.
- Rostek, F., G. Ruhland, F. C. Bassinot, P. J. Muller, L. D. Labeyrie, Y. Lancelot, and E. Bard, Reconstructing sea surface temperature and salinity using  $\delta^{18}O$  and alkenone records, *Nature*, 364, 319–321, 1993.
- Sachs, J. P., and S. J. Lehman, Subtropical North Atlantic temperatures 60,000 to 30,000 years ago, *Science*, 286, 756–759, 1999.
- Sicko-Goad, L., M. S. Simmons, D. Lazinsky, and J. Hall, Effect of light cycle on diatom fatty acid composition and quantitative morphology, *J. Phycol.*, 24, 1–7, 1988.
- Strickland, J. D. H., and T. R. Parsons, *A Practical Handbook of Seawater Analysis*, 2nd ed., *Bull. Fish. Res. Board Can.*, 167, 310 pp., 1972.
- van der Wal, P., J. J. M. Leunissen-Bijvelt, and A. J. Verkleij, Ultrastructure of the membranous layers enveloping the cells of the coccolithophorid *Emiliania huxleyi*, *J. Ultrastruct. Res.*, 91, 24–29, 1985.
- Verardo, D. J., P. N. Froelich, and A. McIntyre, Determination of organic carbon and nitrogen in marine sediments using the Carlo Erba NA-1500 analyzer, *Deep Sea Res., Part I*, 37, 157–165, 1990.
- Volkman, J. K., D. A. Everitt, and D. I. Allen, Some analyses of lipid classes in marine organisms, sediments and seawater using thin-layer chromatography-flame ionization detection, *J. Chromatogr.*, 356, 147–162, 1986.

F. G. Prahl and M. A. Sparrow, College of Oceanic and Atmospheric Sciences, Oregon State University, Oceanography Administration Building 104, Corvallis, OR 97331-5503, USA. (fpahl@coas.oregonstate.edu; msparrow@coas.oregonstate.edu)

G. V. Wolfe, Department of Biological Sciences, California State University, Chico, CA 95929-0515, USA. (gwolfe2@csuchico.edu)